

# The climate change variations in the northern Greater Khingan Mountains during the past centuries

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**Abstract:** The Greater Khingan Mountains (Daxinganling) are China's important ecological protective screen and also the region most sensitive to climate changes. To gain an in-depth understanding and reveal the climate change characteristic in this high-latitude, cold and data-insufficient region is of great importance to maintaining ecological safety and corresponding to global climate changes. In this article, the annual average temperature, precipitation and sunshine duration series were firstly constructed using tree-ring data and the meteorological observation data. Then, using the climate tendency rate method, moving-t-testing method, Yamamoto method and wavelet analysis method, we have investigated the climate changes in the region during the past 307 years. Results indicate that, since 1707, the annual average temperature increased significantly, the precipitation increased slightly and the sunshine duration decreased, with the tendency rates of  $0.06^{\circ}\text{C}/10\text{a}$ ,  $0.79\text{ mm}/10\text{a}$  and  $-5.15\text{ h}/10\text{a}$ , respectively ( $P \leq 0.01$ ). Since the 21st century, the period with the greatest increase of the annual average temperature (also with the greatest increase of precipitation) corresponds to the period with greatest decrease of sunshine duration. Three sudden changes of the annual average temperature and sunshine duration occurred in this period while two sudden changes of precipitation occurred. The strong sudden-change years of precipitation and sunshine duration are basically consistent with the sudden-change years of annual average temperature, suggesting that in the mid-1860s, the climatic sudden change or transition really existed in this region. In the time domain, the climatic series of this region exhibit obvious local variation characteristics. The annual average temperature and sunshine duration exhibit the periodic variations of 25 years while the precipitation exhibits a periodic variation of 20 years. Based on these periodic characteristics, one can infer that in the period from 2013 to 2030, the temperature will be at a high-temperature stage, the precipitation will be at an abundant-precipitation stage and the sunshine duration will be at an less-sunshine stage. In terms of spatial distribution, the leading distribution type of the annual average temperature in this region shows integrity, i.e., it is easily higher or lower in the whole region; and the second distribution type is more (or less) in the southwest parts and less (or more) in the northeast parts. Precipitation and sunshine duration exhibit complex spatial distribution and include four

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spatial distribution types. The present study can provide scientific basis for the security investigation of homeland, ecological and water resources as well as economic development programming in China's northern borders.

**Keywords:** climate change; temporal and spatial characteristics; northern Greater Khingan Mountains; sudden change test; period analysis

Forest has a close and complex relationship with climate. Climatic conditions are decisive factors for forest attributes. In September, 2013, Intergovernmental Panel on Climate Change issued the fifth assessment report on climate changes, which pointed out that global climate changes have imposed significant effects on many natural ecosystems on the earth. Especially, the sensitivity of forest ecosystems to global climate changes has led to the accelerated degradation of forest ecosystems on a global scale (IPCC, 2013). Accordingly, the mechanism of interaction between climate changes and forest ecosystems on multiple scales from different aspects is not only a core problem attracting the attention of ecologists and climatologists, but also the focus and a frontier question sharing the concerns of international community currently (Xie *et al.*, 2009; Yang *et al.*, 2012). With regard to the studies on climate changes based on tree ring, the foreign researches have conducted the tree-ring-based climate reconstruction for a long time and many studies were focused on the climate reconstruction based on tree-ring network in many arid and cold regions of North America (Briffa *et al.*, 1998; Cook *et al.*, 2002; D'Arrigo *et al.*, 2005; Yadav *et al.*, 2002). To construct a perfect tree-ring network, more than a hundred of regional tree-ring samples are required, in which temperature, precipitation, drought index and North Pacific Index should be adopted as the factors. Many satisfactory results have already been achieved in climatic sudden change analysis. The investigations on the climate changes in high-latitude regions of China over a century have started late, but with a high starting point. In these studies, the researches not only paid attention to the variation analysis from single meteorological factor (Cai *et al.*, 2010; Gou *et al.*, 2013; Shi *et al.*, 2010), but also explored the characteristics and effects of regional climate changes (Lu *et al.*, 2014; Sun *et al.*, 2012; Wang *et al.*, 2012; Yu *et al.*, 2005). These achievements have laid solid theoretical basis and technical supports for revealing the evolvement rules of climate changes in the Greater Khingan Mountains (GKM), a high-latitude cold region.

The Greater Khingan (GKM) region is a first-grade National Nature Ecological Reserve, which is located at the southern margin of the permafrost of Eurasia (Wu *et al.*, 2014). The GKM is in the alpine region, an environmental sensitive area in Northeast Asia. This region is abundant in forest resources, including China's only primeval cold-temperature bright coniferous forest with the largest area and best preserve, with the constructive species of *Larix gmelinii* and Mongolian pine. Compared with the other forest regions in China, the GKM exhibits a series of unique features in terms of soil, disturbance regime, climate and vegetation. Few systematic studies were conducted on the climate changes in the northern GKM; moreover, due to the limitations in high-latitude data density and series length, the research achievements in regional climate changes lack of systematicness and comprehensiveness. Therefore, using the reconstructed data series of annual average temperature, precipitation and sunshine duration based on tree-ring width during the period from 1707 to 2013 in the northern GKM, we conducted a systematic investigation on temporal and spatial

characteristics of the climate changes in the northern GKM, so as to further determine the rules of climate changes and identify the change characteristics. The present study aims to gaining an in-depth insight into the climate changes in this alpine region, which can not only lay solid foundations for the climate change investigations in the whole high-latitude region, but also provide significant scientific and application values for the homeland, ecological and water resource security as well as economic development programming in China's northern borders.

## 1 Materials and method

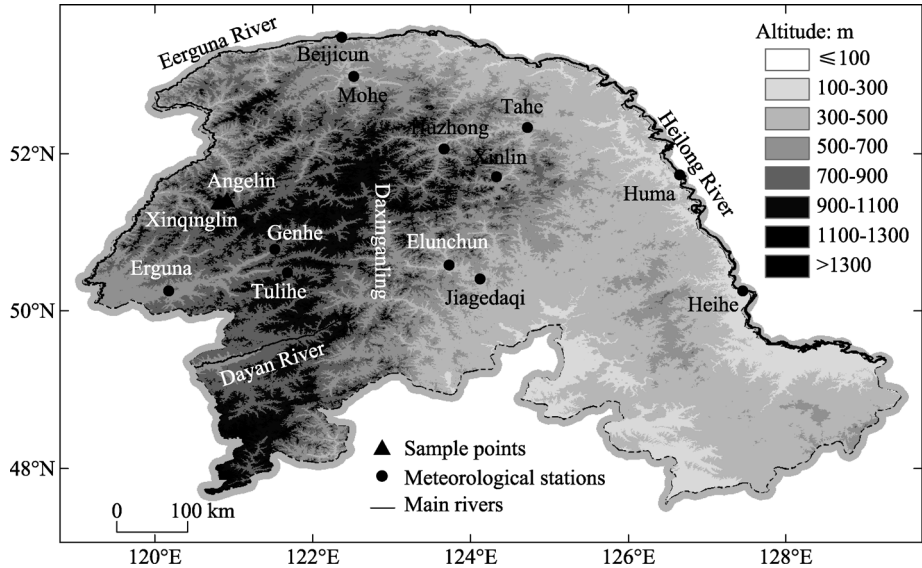
### 1.1 Data acquisition and a general introduction to the study region

The northern GKM was selected as a case in the present study, with the latitude ranging from 50°10' to 53°30'N and the longitude ranging from 119°40' to 127°22'E. It is bounded on the north and northeast by the Heilongjiang River while on the west and northwest by the Erguna River, bordering to East Siberia and Far East in Russia across the two rivers each; the southeast extends to the line along Heihe City and Nenjiang County and is connected to the Lesser Khingan Mountains (Range). The study region covers an area of approximately  $16.22 \times 10^4 \text{ km}^2$  (exceeding the area of Liaoning Province). In terms of vegetation regionalization, it is an independent cold-temperate coniferous forest zone and characterized by a cold-temperate continental monsoon climate (Zhou, 1991). From south (east) to north (west), the annual temperature decreases from  $-1-0^\circ\text{C}$  to  $-5-4^\circ\text{C}$ , the annual temperature range increases from  $35^\circ\text{C}$  to  $52^\circ\text{C}$ , and the annual precipitation decreases from 500–700 mm to below 200 mm (He *et al.*, 2013). The soils are mainly dark brown coniferous forest soils, and the vegetation cover is the Khingan flora, in which the Khingan larch trees occupy a large proportion in tree species, followed by Mongolian pine and dragon spruce.

In the present study, the climatological data of the northern GKM consists of the climatic factors inverted by the tree-ring width data from 1707 to 2007 and the climatological observation data from 1961 to 2013. After being tested, the historical climatological data series are from 1707 to 2013 (lasting 307 years), with the climatic factors of annual average temperature, precipitation and sunshine duration. It should be noted that the data of all these climatic factors were calculated by averaging the data from April in the current year to March in the next year. According to requirements in the present study, these data series include three sections (Li *et al.*, 2011; Li *et al.*, 2014): specifically, from 1707 to 1960, since no instrument-measured data in this region, the climatological data were acquired by inverting the observation data from Erguna Municipal Meteorological Station (50°15'0"N, 120°10'48"E, with a height of 581 m) and Genhe Meteorological Station (50°46'48"N, 121°31'12"E, with a height of 779 m) through the constructed regression equations (these two meteorological stations are at the nearest distance from Angelin Forest Farm (51°25'12"N, 120°54'51"E) and Xinqing Forest Farm (51°23'3"N, 120°48'38"E) and have the same underlying surfaces and the tree-ring width data were collected in these two forest farms); from 1961 to 2007, the climatological data were calculated by averaging the data in Angelin and Xinqing tree-ring width inversion stations and the observation data of 11 meteorological stations (the station density was added); from 2008 to 2013, the climatological data were the average values of the observation data collected from 11 meteorological sta-

tions (the tree-ring width data were collected in 2007). Since the Beijicun Station was built later and the data series are incomplete, the data were deleted (only used for matching).

Based on the standard tree-ring chronology and meteorological data, we have investigated the meteorological data during the period from 1707 to 2007 using regression method (Qu *et al.*, 2016). As stated in the previous research results, the constructed quantitative relationship between tree-ring width and meteorological factors could pass the credibility test. Moreover, the obtained 307-year meteorological data series could pass the consistency check (see details in Qu *et al.*, 2016). More importantly, compared with the specific meteorological data, we are more concerned with the data series length, the density of the acquired data, the fluctuation changes of various meteorological factors, and the characteristics of regional climate changes. Therefore, the processing meteorological data series in this article can meet the technical requirements that the tree ring can be used for replacing meteorological data, and the meteorological series have certain reliability in the applications of present study. Figure 1 illustrates the distributions of tree-ring sample points and meteorological stations.



**Figure 1** Illustration of the geological positions of the northern Greater Khingan Mountains and the distributions of meteorological stations and tree-ring sampling points

**1.2 Method**

**1.2.1 Wavelet analysis**

The descriptions of the time series’ pattern evolution characteristics and the position in each period at each moment using wavelet theory can enhance the short-time Fourier transform’s ability in dealing with practical problems and has become a new method of analyzing the series’ time-varying characteristics.

Assuming a wavelet function that satisfies a certain condition, the wavelet transform of the time series can be described as:

$$W_f(a,b)=|a|^{-\frac{1}{2}}\int_Rf(t)\overline{\Psi}\left(\frac{t-b}{a}\right)dt \tag{1}$$

where  $W_f(a, b)$  denotes wavelet transform or wavelet coefficient, and varies with the variations of  $a$  and  $b$  ( $a$  denotes a frequency parameter and reflects the length of wavelet period,  $a > 0$ , while  $b$  denotes a time parameter and reflects the wave's translation over time). The wavelet coefficient  $W_f(a, b)$  on different time scales (periods, with the unit of year, the same below) can reflect the system's variation characteristics on different time scales. The turning point between positive and negative wave coefficients corresponds to the sudden change point. The greater the wavelet coefficient, the more significant the variations on this time scale (Ge *et al.*, 2013; Guo *et al.*, 2013; Zhao *et al.*, 2012).

Mexican Hat Function can be expressed as:

$$\Psi(t) = (1 - t^2)e^{(-t^2/2)} \quad (2)$$

The wavelet variance test was conducted for determining the main periods of several data series of annual average temperature, precipitation and sunshine duration:

$$W_p(a) = \int_{-\infty}^{\infty} |W_f(a, b)|^2 db \quad (3)$$

where  $W_p(a)$  denotes the wavelet variance. On a certain time scale, wavelet variance reflects the periodic fluctuation intensity of the time series on this scale. The scale corresponding to the peak value thus can be regarded as the main time scale of this series, i.e., the main period.

### 1.2.2 Moving t-test (MTT)

The principle of moving t-test (MTT) is detailedly described below. The sudden change can be tested according to the fact whether the difference exists between two sets of sample average values. For a series  $x$  with the total sample size of  $n$ , a certain moment can be artificially set as the datum point, and two sub-series before and after this datum point with the length of  $n_1$  and  $n_2$  (generally  $n_1 = n_2$ ) are selected for moving calculation. Then, a statistic series of  $t$  can be acquired. The critical value  $t_\alpha$  can be determined by the given quantitative level. If  $|t| < t_\alpha$ , the two sub-series before and after this datum point are considered to be with no significant difference; otherwise, it can be assumed that the sudden change occurs at the datum point. MTT method can be expressed as:

$$t = \frac{\bar{x}_1 - \bar{x}_2}{s \cdot \sqrt{\frac{1}{n_1} + \frac{1}{n_2}}}, \quad s = \sqrt{\frac{n_1 s_1^2 + n_2 s_2^2}{n_1 + n_2 - 2}} \quad (4)$$

where  $s_1$  and  $s_2$  are the standard deviations of two sub-series. In this article,  $n=307$  years,  $n_1$  and  $n_2$  were set as 20 years, and  $\alpha=0.01$  was adopted as the significance level of sudden change judgment, i.e., in the continuous years at different periods passed the test of significance level, the year with the greatest value of  $t$  was selected as the strongest sudden-change year.

### 1.2.3 Yamamoto method (YAMA)

Using MTT, since the moving step size was artificially set (Sun *et al.*, 2005; Zhang *et al.*, 2015), the sudden change point may drift. In the present study, the sudden change analysis

was firstly conducted using MTT; the test was conducted on the detected sudden change point (if exists) using Yamamoto method (YAMA); if the results using two methods are consistent or approach with each other, the sudden change year can be identified.

Using YAMA method, the significant difference of the average values at different times can be tested by the signal to noise ratio ( $S_{NR}$ ):

$$R_{SN} = \frac{|\bar{x}_1 - \bar{x}_2|}{S_1 + S_2} \quad (5)$$

where  $\bar{x}_1, \bar{x}_2, S_1$  and  $S_2$  denote the average values and variances at two different times.  $n_1$  and  $n_2$  of the time periods for comparison can be set in accordance with requirements, whose values affect the significance level of  $R_{SN}$ . Generally, for a continuous random variable, the subsections are set uniformly, i.e.,  $n=n_1=n_2$ . In the present study,  $n$  was set as 20, if  $R_{SN}>0.60$ , it exceeds the significance level of  $\alpha=0.01$ , and the point can be regarded as the sudden change point; if  $R_{SN}>0.79$ , the point can be regarded as the strong sudden change point.

### 1.2.4 Climate tendency rate method

To calculate the climate tendency rate of the climatic series in the northern GKM (Zhao *et al.*, 2008), the year  $t$  was adopted as the time factor and the climatic factor  $x$  was adopted as the simulation objection; then the linear regression equation between  $x$  and  $t$  was constructed as:  $x(t)=c+bt$ , in which  $c$  and  $b$  are the coefficients to be confirmed. Specifically,  $b$  denotes the tendency of climatic factors. If  $b>0$ , the climatic factors exhibit rising trends; if  $b\leq 0$ , the climatic factors exhibit downtrends. Therefore,  $b\times 10$  is referred to as the climate tendency rate, with the unit of the ratio of the unit of climatic factors to 10 years.

### 1.2.5 Empirical orthogonal function (EOF) analysis

EOF analysis focuses on the matrix data's structural characteristics and aims to extract the primary characteristic quantity of data. It is a common temporal and spatial characteristic distribution method. The annual average temperature, precipitation and sunshine duration were described as matrices:  $X_{m\times n}$  ( $m=1, 2, \dots, 13; n=1, 2, \dots, 43$ ). After standardization processing, these matrices can be decomposed into several new matrices (Huang *et al.*, 2013):

$$N_{m\times n} = V_{m\times p} T_{p\times n} \quad (6)$$

where  $V_{m\times p}$  denotes the spatial characteristic vector and  $T_{p\times n}$  denotes the time coefficient. When  $p=n=43$ , the information of all the meteorological stations can be completely described. It means that the primary variation characteristics of climatic factors can be roughly expressed using the first several characteristic vectors.

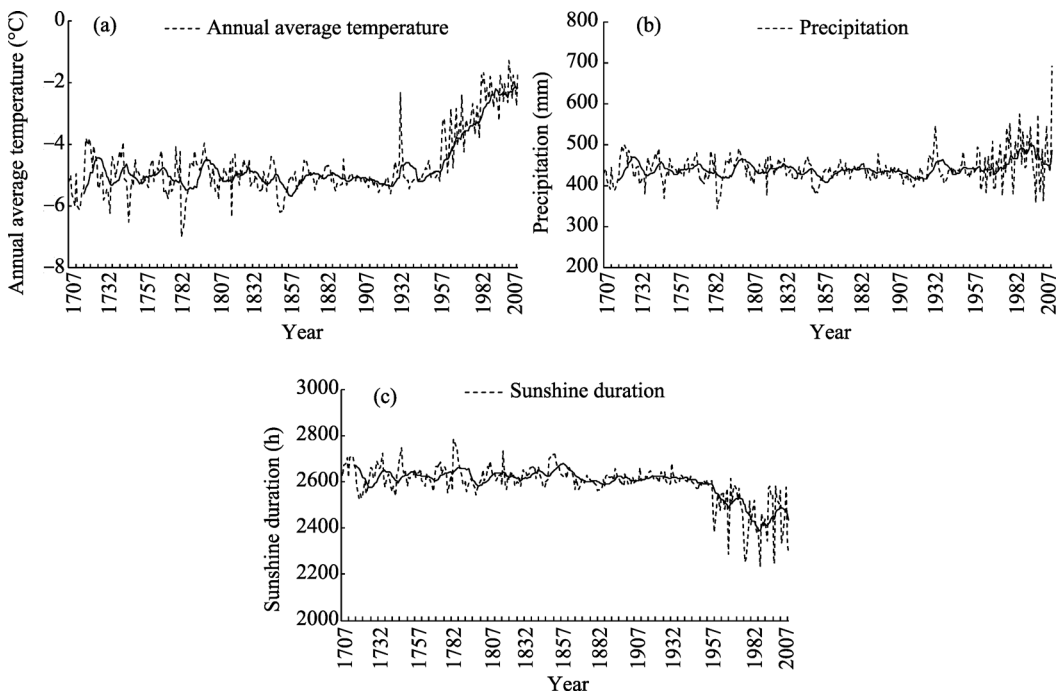
## 2 Results and analysis

### 2.1 Interannual and interdecadal variation characteristics of climate changes

Used climate tend rate, MTT, YAMA and Mexican Hat wavelet function method, the changing trend, abrupt change test and periodic change of climate series in northern GKM were analyzed.

### 2.1.1 Tendency analysis of climate series

Figure 2 displays the climate series of the northern GKM from 1707 to 2013 and the variation curves after the moving average with the moving step size of 11 years. The climate tendency rate was calculated by the first-order linear tendency equation. As shown in Figure 2a, the annual temperature varies between  $-6.99$  and  $-1.28^{\circ}\text{C}$ , with a climate tendency rate of  $0.06^{\circ}\text{C}/10\text{a}$  ( $P \leq 0.01$ ). Before the 20th century, the variation is slight, and the annual average temperature is approximately  $-5.08^{\circ}\text{C}$ . After the 20th century, the annual average temperature increases gradually in the fluctuation; especially, a leap temperature increase appeared in the 1930s. Then the increase of the annual average temperature is smooth again, and this region entered into a continued warming period since the 1950s. During the period from 1951 to 2013, the annual average temperature is approximately  $-3.22^{\circ}\text{C}$ , which exceeds by  $1.86^{\circ}\text{C}$  compared with the data in the 18th and 19th centuries. Moreover, this region is still in a warming stage. As shown in Figure 2b, the precipitation varies between 344.86 and 690.16 mm, with a climate tendency rate of  $0.79 \text{ mm}/10\text{a}$  ( $P \leq 0.01$ ). The precipitation almost exhibits no variations and only increases slightly. During the period from 1707 to the end of the 19th century, the precipitation exhibits slight fluctuation changes, with the average precipitation of 436.71 mm. In the early 20th century, a low-precipitation-value period occurs, with the duration time less than 25 years. By the 1930s, the precipitation increases slightly and the fluctuation amplitude of precipitation increases significantly by the 1980s. During the period from 1981 to 2013, the annual average precipitation is 480.46 mm, exceeding by 43.74 mm compared with the value in the early 20th century. Moreover, the precipitation is in a slow increasing stage. As shown in



**Figure 2** Interannual variation tendency of the climatic series in the northern Greater Khingan Mountains during the period from 1707 to 2013 (a. annual average temperature; b. precipitation; c. sunshine duration, the same below)

Figure 2c, the sunshine duration in the northern GKM is 2233.36–2783.72 h, with a climate tendency rate of  $-5.15 \text{ h/10a}$  ( $P \leq 0.01$ ). Before the 1950s, the sunshine duration exhibits no obvious variation tendency; after the 1950s, it decreases significantly, which at least shows that the sprinkles and hazy weathers increase gradually.

Table 1 lists the variation of climate series in the northern GKM over different centuries from 1707 to 2013. One can observe that, the annual average temperature exhibits no distinct variations in the 18th and 19th centuries; but increases significantly in the 20th century. The annual average temperature in the 20th century is  $0.83^\circ\text{C}$  higher than the value in the 19th century, with the warming amplitude of  $0.32^\circ\text{C/10a}$ . The results also passed the significance tests. The precipitation in this region decreases slightly in the 19th century, which is 3.73 mm lower than the value in the 18th century. After 1900, the precipitation increases gradually, with a climate tendency rate of  $6.04 \text{ mm/10a}$  ( $P \leq 0.01$ ). During the period from 2000 to 2013, the precipitation is the greatest. On the whole, the sunshine duration decreases in this region. The sunshine duration from 1800 to 1999 is 2.75 h lower than the value in the 18th century; after 1901, the sunshine duration decreases significantly, with a climate tendency rate of  $-23.27 \text{ h/10a}$  ( $P \leq 0.01$ ). From 2000 to 2013, the sunshine duration decreases most remarkably, and the sunshine duration during this period is reduced by 100.14 h compared with the value in the 20th century.

**Table 1** Variation of climate series in the northern Greater Khingan Mountains from 1707 to 2013

Year	Annual average temperature		Precipitation		Sunshine duration	
	Average value/ $^\circ\text{C}$	Climate tendency rate/ $^\circ\text{C/10a}$	Average value/mm	Climate tendency rate/mm/10a	Average value/h	Climate tendency rate/h/10a
1707–1799	−5.04	0.02	438.64	0.85	2626.87	−1.44
1800–1899	−5.11	−0.01	434.91	−0.33	2624.13	−2.80*
1900–1999	−4.28	0.32**	446.50	6.04**	2551.71	−23.27**
2000–2013	−2.15	−0.01	470.44	116.57	2451.57	−116.22

Note: \* denotes the results passed the significant test at the level of  $P = 0.05$ ; \*\* denotes the results passed the significant test at the level of  $P = 0.01$

**2.1.2 Sudden change detection of climatic series**

Using MTT and YAMA method, the sudden change detection was performed on the annual average temperature, precipitation and sunshine duration series during the period from 1707 to 2013 in the northern GKM, and the reliable sudden change years were determined through comparisons. Table 2 lists the results of years that passed the significance test at the level of  $P \leq 0.01$ . One can conclude that the annual average temperature exhibits sudden changes in 1866, 1960 and 1987, in which the sudden change in 1960 is most intensive; the precipitation exhibits sudden changes in 1867 and 1927, in which the sudden change in 1867 is most intensive; the sunshine duration exhibits sudden changes in 1866, 1960 and 1978, in which the sudden change in 1866 is most intensive. The years with most strong sudden changes of precipitation and sunshine duration are basically consistent with the years with sudden changes of annual average temperature, indicating that the climate in this region really underwent sudden changes or transitions in the mid-1860s. The sudden change directions were further determined using MTT analysis. It can be observed that annual average temperature



and sunshine duration share the same sudden change points but different sudden change directions. The annual average temperature and precipitation varied suddenly from less to more while the sunshine duration varied suddenly from more to less, which are in good consistency with the variation tendencies of the above climatic series.

**Table 2** Sudden change years of the climatic series in the northern Greater Khingan Mountains from 1707 to 2013

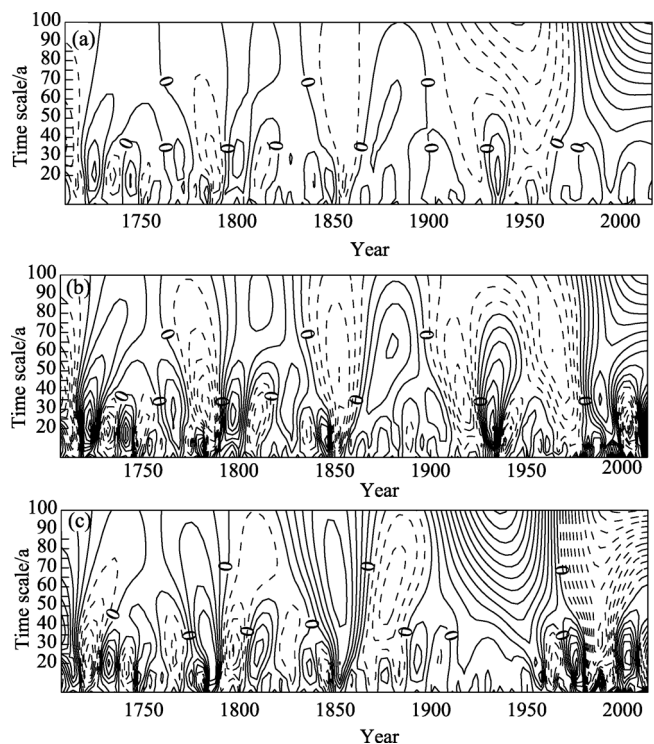
Factors	Methods	Sudden change years									
Annual average temperature	MTT	1771	1788	1846	1866	1960	1987				
	YAMA	–	–	–	1866	1960	1987				
Precipitation	MTT	1771	1788	1847	1867	1900	1908	1927	1974	1979	
	YAMA	–	–	–	1867	–	–	1927	–	–	
Sunshine duration	MTT	1771	1788	1866	1887	1934	1960	1978			
	YAMA	–	–	1866	–	–	1960	1978			

### 2.1.3 Periodic variation of climatic series

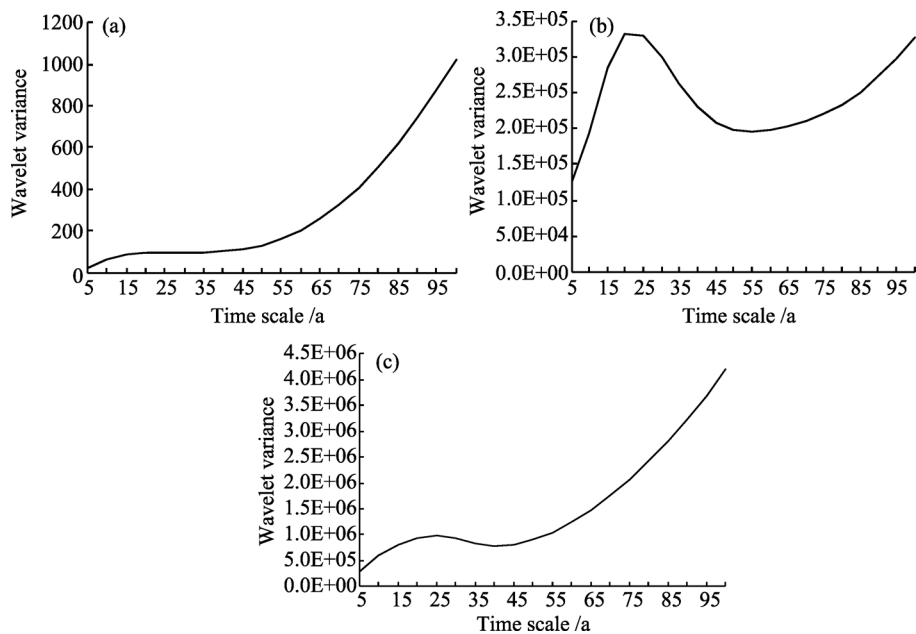
Figure 3 displays the Mexican Hat wavelet analysis results of the climatic series from 1707 to 2013, in which the negative isolines were denoted by the dash lines and represent that the factor values are comparatively lower, and the positive or zero isolines were denoted by the solid lines and represent that the factor values are comparatively greater. As shown in Figure 3a, the periodic oscillation of annual average temperature is intensive on a time scale below 10 years, exhibiting no obvious rules; as the time scale increases, the oscillations show obvious periodic rules on the time scale of 20–25 years, and the fluctuation variations on the calculated time domain can be observed, i.e., the annual average temperatures are sometimes lower and sometime higher and totally 8 alternations of warming and cooling occurred in this region. Moreover, the isolines of annual average temperature have still not been closed by 2013, suggesting that this region is in a warming period for some time to come. As shown in Figure 3b, the periodic oscillation of precipitation is also intensive on a time scale below 10 years, exhibiting no obvious rules; with the increase of time scale, the obvious rules can be observed in the periodic oscillation on the time scale of 20–25 years and totally 9 alternations of drying and wetting occurred in this region. Similarly, isolines of precipitation have still not been closed by 2013, suggesting that the precipitation in this region will be slightly more for some time in the future. As shown in Figure 3c, on the time scale of 25–30 years, the sunshine duration in this region shows significant periodic oscillations and totally 10 alternations of sunny and rainy can be observed. Until 2013, the isolines of sunshine duration have still not been closed, suggesting that the sunshine duration will be slightly less for some time in the future.

Figure 4 displays the period analysis on the climatic series in the northern GKM during the period from 1707 to 2013 using the variance of wavelet coefficient. One can observe that, from 1707 to 2013, the wavelet variance of annual average temperature is 25 years; the wavelet variance of precipitation reaches a peak at 20 years, i.e., the primary period of the precipitation variation is 20 years; the sunshine duration variations are mainly the oscillations with the quasi-period of 25 years. According to the periodic characteristics, one can speculate that the annual average temperature will be in a wavelet period of continued

warming from 2013 to 2030, the precipitation will be in a wavelet period characterized by slightly more precipitation from 2013 to 2029, and the sunshine duration will be in a wavelet period characterized by slightly less duration from 2013 to 2034. In other words, during the



**Figure 3** Wavelet analysis of climatic series in the northern Greater Khingan Mountains

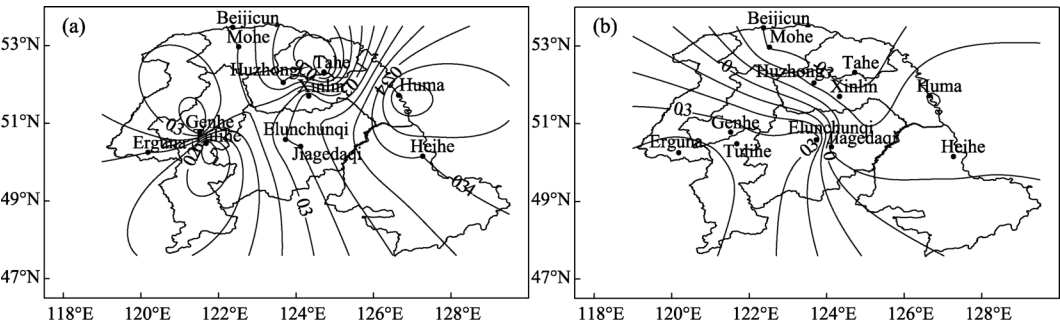


**Figure 4** Wavelet variances of climatic series in the northern Greater Khingan Mountains

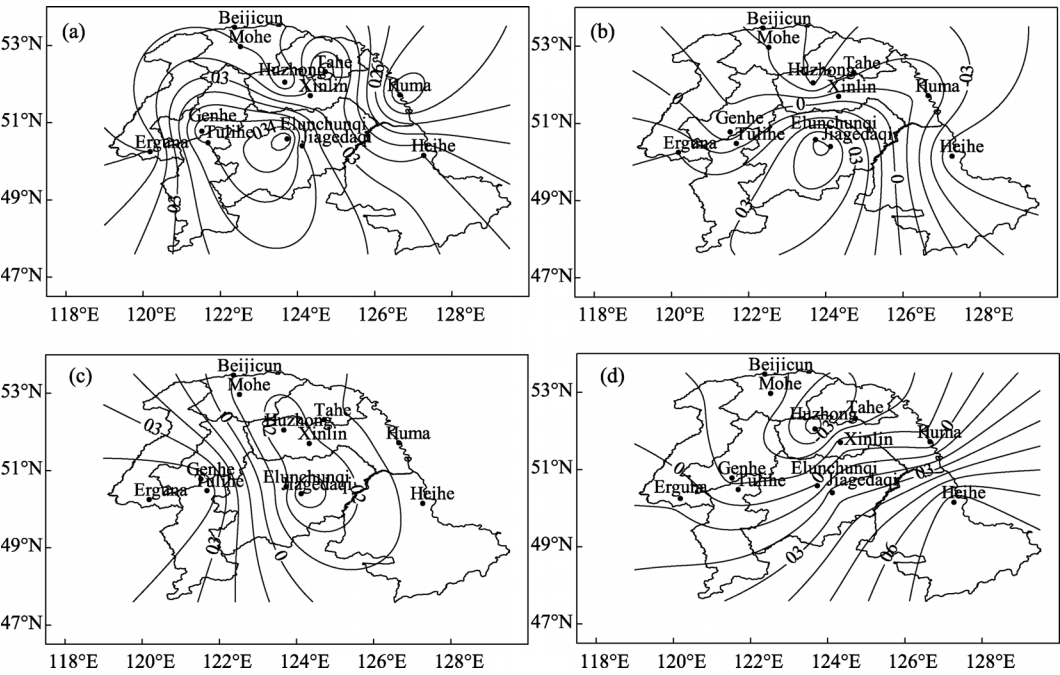
period from 2013 to 2030, the annual average temperature will be in a slightly-higher period, the precipitation will be in a slightly-more period and the sunshine duration will be in a slightly-less period.

## 2.2 Spatial analysis of climate changes

To analyze the spatial climate changes in the northern GKM, standardization processing and EOF analysis were conducted on the climatic series in this region, as the results presented in Figures 5–7.

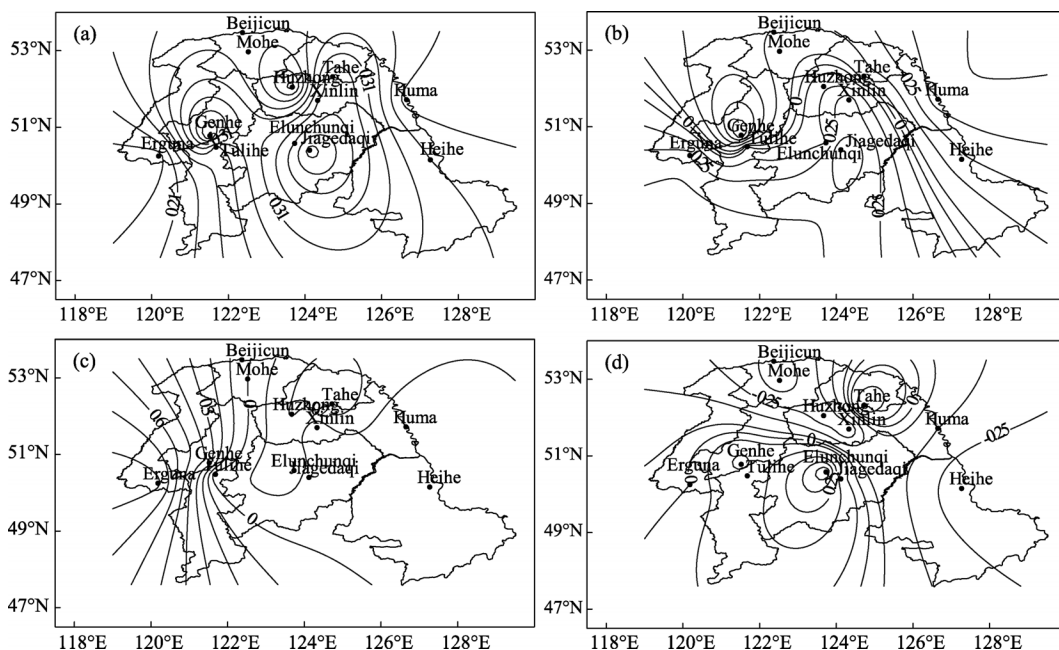


**Figure 5** Eigenvector fields of annual average temperature in the northern Greater Khingan Mountains (a. The first eigenvector field; b. The second eigenvector field)



**Figure 6** Eigenvector fields of precipitation in the northern Greater Khingan Mountains (a. The first eigenvector field; b. The second eigenvector field; c. The third eigenvector field; d. The fourth eigenvector field, the same below)

As shown in Figure 5, after EOF expansion on the annual average temperature series, the variance contribution rate of the first eigenvector can reach up to 82.8%; the accumulated



**Figure 7** Eigenvector fields of sunshine duration in the northern Greater Khingan Mountains

variance contribution rate of the first two eigenvectors is as high as 93.3%, suggesting that the first two eigenvectors can represent the variation of annual average temperature in this region. The typical field corresponding to the first eigenvector is positive and shows the consistency of annual average temperature variation in this region. As displayed in this typical field, the high-value distribution is concentrated along the Huma-Heihe line, and the eigenvectors values decrease gradually to the west and north; the secondary- high-value distribution is at the center of Genhe. The results demonstrate that, in these two regions of the northern GKM, the annual average temperature variation rates are relatively large; these two regions are also the sensitive regions to temperature changes. The typical field corresponding to the second eigenvector shows the variation characteristics that the southwest part is positive and the northeast part is negative. Zero-value lines are extended towards the north-west direction through Jiagedaqi and Elunchunqi, which represent the second temperature types in the study region. This type shows that the annual average temperature in this region shows the pattern characteristics of higher in the southwest part and lower in the northeast part or lower in the southwest part and higher in the northeast part.

Figure 6 displays the EOF analysis results of the precipitation series in the northern GKM. The accumulated variance contribution rate of the first four eigenvectors can reach up to 83.2%, suggesting that the first four eigenvectors can roughly represent the precipitation in the study region. These four eigenvectors were selected as the four basic distribution types in the study region, and thereby the precipitation in the study region can be classified into four types, namely, more (or less) in the whole region; more (or less) in the central region and less (or more) in the northeastern region; more (or less) in the eastern region and less (or more) in the western region; and more (or less) in the southern region and less (or more) in the northern region. The contribution rate of the first eigenvector to the total variance is up

to 55.9%, and the typical field corresponding to the first eigenvector can be regarded as the most important typical field reflecting the precipitation variations in the study region. All this typical field is positive, i.e., the whole region is rainy or rainless and can be classified as the first distribution type. The high-value distribution is at the center of Elunchunqi, and the value decreases gradually to the north and east, reflecting that the region's precipitation varies greatly. The contribution rate of the second eigenvector to the total variance is up to 11.2%, and zero-value isolines pass through Yilehuli Mountain and GKM. In terms of spatial distribution, the positive value region is located at the center of Elunchunqi that are surrounded by two mountains while the negative value region is located in the eastern Yilehuli Mountain and the northern GKM. The third eigenvector corresponds to the type of more (or less) in the eastern region and less (or more) in the western region, with the variance contribution rate of 8.8%. In terms of spatial distribution, the negative values in the eastern regions increase gradually towards the west, with the negative-value center in the Elunchunqi, i.e., the spatial pattern is characterized by the reverse variations between west and east. The contribution rate of the fourth eigenvector to the total variance is up to 7.4%. An opposed phase between north and south can be observed. A negative-value center exists in the north of Huzhong while the positive-value center is located in the south of Heihe, reflecting the fourth type of more (or less) in the southern region and less (or more) in the northern region. In other words, in this region, the precipitation is characterized by the reverse variations between south and north.

Figure 7 displays the EOF analysis results of the sunshine duration series in the northern GKM. The accumulated variance contribution rate of the first four eigenvectors can reach up to 85.0%, i.e., the first four eigenvectors can reasonably represent the spatial abnormal distribution of sunshine duration in the study region. The spatial distribution of the first eigenvector is characterized by the positive values in the whole region and in-phase distribution, suggesting that the spatial distribution of sunshine duration in the study region is mainly the consistency-type, i.e., the sunshine duration is easily more or less in the whole region. The maximum-value center is located at the center of Elunchunqi and the second high-value center is located at the center of Genhe. Since the contribution rate the first eigenvector to the total variance is up to 54.6%, more or less sunshine duration in the whole region is the leading spatial distribution type in the study region. The contribution rate of the second eigenvector to the total variance is up to 14.0%, suggesting that the variation type corresponding to the second eigenvector is also a typical spatial distribution of sunshine duration variation. This distribution type uses the Yilehuli Mountain and the GKM as the boundary. The values in the southern parts are positive while the values in the northern parts are negative. The positive-value center is located at the center of Elunchunqi, while the negative-value center is located at the center of Genhe. This spatial distribution is mainly affected by the geological characteristics. The contribution rate of the third eigenvector to the total variance is up to 9.3%. On the whole, the spatial distribution took the Mohe-Tulihe line as the boundary; the values in the eastern parts are negative while the values in the western parts are positive. This is a typical more (or less) in the eastern region and less (or more) in the western region. The greatest positive-value center is in the southern Eerguna while the greatest negative-value center is located in Tahe. The fourth eigenvector field varied greatly from the first three eigenvector fields. In terms of spatial distribution, the values in the mid-

dle parts are positive while the values in the eastern and western parts are negative. The positive-value center is located at the center of Jiagedaqi and Tahe. The contribution rate of the fourth eigenvector to the total variance is up to 7.1%, i.e., the fourth eigenvector reflects the fourth distribution type of sunshine duration in this study region.

2.3 Comparisons with the other research results

Table 3 lists the comparisons between the present research results and the results for Northeast China during the period from 1951 to 2007 as described by Fu *et al.* (2009). One can observe that, during the past 60 years, the warming tendency in the northern GKM is consistent with the warming tendency in Northeast China, with the annual average temperature tendency rates of 0.54°C/10a and 0.60°C/10a, respectively; the precipitations in the northern GKM and Northeast China both exhibit no obvious variation, with the precipitation tendency rates of 5.67 mm/10a and –0.27 mm/10a, but the variation tendencies are different during this period. Temperature and precipitation show a slight difference in spatial distributions, with almost identical distribution characteristics. By comparing the results from 1989 to 2007 with the results from 1971 to 1988, one can observe that spatial distributions of annual average temperature variations are basically identical, with the temperature increasing amplitudes of 0.67–1.73°C and 0.98°C, respectively. As described by Fu *et al.* (2009), the temperature increasing amplitudes are slightly lower except the results in the northern GKM, i.e., certain differences can be observed. The results fit well the tendency of global warming, which is due to the multiplication of greenhouse gases and the variation of solar radiation.

**Table 3** Comparisons of the annual average temperature and precipitation between the present study and the previous results in the northern Greater Khingan Mountains during the period from 1951 to 2007

Areas	Meteorological Factors	Comparisons during the periods from 1989 to 2007 and from 1971 to 1988, respectively	Distribution areas	Climate tendency rates from 1951 to 2007 (°C, mm/10a )
Northeast China (from the literature)	Annual average temperature variation	0.67–1.73	Hulunbuir Pasture Land, Greater Khingan Mountains, Lesser Khingan Mountains, Songnen Plain, Liaoxi Mountain Land, Liaodong Peninsula, Sanjiang Plain, Changbai Mountains	0.60
	Annual precipitation variation	5.40–67.27	Yilehuli Mountains, northern Greater Khingan Mountains, Lesser Khingan Mountains, northern Changbai Mountains, western Hulunbuir Pasture Land, eastern Inner Mongolian Plateau	–2.70
Northern Greater Khingan Mountains (from this article)	Annual average temperature variation	0.98	Northern Greater Khingan Mountains (50°10′–53°30′N, 119°40′–127°22′E)	0.54
	Annual precipitation variation	13.26	Northern Greater Khingan Mountains (50°10′–53°30′N, 119°40′–127°22′E)	5.67

On the one hand, the northern GKM are mainly primeval forests, with sparse population. The temperature variation is tightly correlated with the topographic changes, and the temperature increasing amplitude decreases gradually with the rising of terrain. On the other hand, outside the northern GKM, the latitude is relatively low, with the vegetation of forests, grasslands and farmlands. The artificial heating and urbanization on account of dense population may also contribute to the warming in this region. The annual precipitation varies slightly in different regions, with the maximum annual precipitation increasing amplitude of 3.5 mm, which can almost be neglected. One can conclude that, in terms of the variations during two periods, the spatial variation amplitude of precipitation decreases gradually from high-latitude region to low-latitude region. This is due to the fact that, in the low-latitude regions, the artificial heating and urbanization have strong effects, the vegetation becomes sparser, and the landform transits from mountain lands to plain, i.e., the Northeast China Plain was formed. This may affect the general atmospheric circulation and the activities of cyclone. Compared with the results in Fu *et al.* (2009), the increasing amplitude of precipitation in the present study is smaller. This is because that the study regions are different and thus the average values of regional precipitation are different. On the whole, our research conclusions are reliable compared with the results in the other region.

### 3 Conclusions

In conclusion, the climate has been experiencing a significant change in the past 300 years in the northern GKM. There is an increasing trend of temperature. The precipitation increases slightly while the sunshine duration decreases. The variations of meteorological factors show distinctly regional difference.

(1) Decadal variations of climatic factors. In the northern GKM, the annual average temperature increases at a rate of  $0.06^{\circ}\text{C}/10\text{a}$  ( $P \leq 0.01$ ) during the period from 1707 to 2013. Specifically, in the 18th and 19th centuries, the annual average temperature in this region exhibits no obvious variations; in the 20th century, the temperature increases significantly, exceeding by  $0.83^{\circ}\text{C}$  than the value in the 19th century, with the increasing amplitude of  $0.32^{\circ}\text{C}/10\text{a}$  ( $P \leq 0.01$ ). The precipitation increased slightly on the whole, with the climate tendency rate of  $0.79 \text{ mm}/10\text{a}$  ( $P \leq 0.01$ ). The precipitation decreased slightly in the 19th century and increases in the 20th century, with the climate tendency rate of  $6.04 \text{ mm}/10\text{a}$  ( $P \leq 0.01$ ). During the period from 2001 to 2013, the precipitation is the greatest. Overall, the sunshine duration decreases, with the climate tendency rate of  $-5.15 \text{ h}/10\text{a}$  ( $P \leq 0.01$ ). The sunshine duration in the 19th century is reduced by 2.75 h compared with the value in the 18th century; in the 20th century, the sunshine duration decreases significantly, with the climate tendency rate of  $-23.27 \text{ h}/10\text{a}$  ( $P \leq 0.01$ ). From 2001 to 2013, the sunshine duration exhibits a greatest decline, which is 100.14 h smaller than the value in the 20th century.

(2) Abrupt change of climatic factors. The annual average temperature became warmer in 1866, 1960 and 1987, the precipitation increased suddenly in 1867 and 1927, and the sunshine duration decreased suddenly in 1866, 1960 and 1978. The strong sudden-change years of precipitation and sunshine duration basically coincide with the sudden-change years of average temperature, suggesting that in the mid-1860s, sudden change or transition of cli-

mate really occurred in the northern GKM. The climatic factors all exhibit obvious periodic changes and show different alternations with different periods. The annual average temperature and sunshine duration both have the primary periods of 25 years, while the precipitation has a significant period of 20 years.

(3) Spatial characteristics of climatic factors. The annual average temperature is characterized by integrality in the northern GKM. The high-value distribution is concentrated along the line of Huma ( $-0.7^{\circ}\text{C}$ ) and Heihe ( $0.7^{\circ}\text{C}$ ), and the eigenvectors values decrease gradually to the west and north; The secondary-high-value distribution is at the center of Genhe. The results suggest that, the annual average temperature variability is great in these two regions of northern GKM, where the temperature changes are also sensitive. As to the precipitation, the high-value distribution is at the center of Elunchunqi, and the value decreases gradually to the north and east. The spatial distribution of sunshine duration is mainly the consistency-type. The maximum-value center is located at the center of Elunchunqi and the secondary high-value center is located at the center of Genhe.

## 4 Discussion

The previous studies on the climate changes on the time scale over a century indicate that, since the 1950s, the concentrations of carbon dioxide ( $\text{CO}_2$ ) increase gradually under the influence of human activities and the earth has entered into an unprecedented era of climate changes (Stocker *et al.*, 2013). Although the researchers now have accepted the viewpoint that the climate changes on a large time scale (for example, the global average temperature increases), the variation characteristics of various climatic factors on regional scale are still hard to determine (Millar *et al.*, 2007). In other words, climate changes exhibit great uncertainties in different regions within a short time period (Walther *et al.*, 2002), which is resulted from many reasons. In the present study, the annual average temperature in the northern GKM increases at a rate of  $0.06^{\circ}\text{C}/10\text{a}$  ( $P \leq 0.01$ ). Although the present results basically fit well with the latest research results (the annual average temperature increases at a rate of  $(0.04\text{--}0.10^{\circ}\text{C})/10\text{a}$  (Ren *et al.*, 2014), they are different from the previous estimations on China's warming tendency over the past century. On the one hand, it indicates that our reconstructed meteorological data series in the northern GKM have high reliability (Tang *et al.*, 2005); on the other hand, the differences originate from the differences in the study range, the length of climatic series and data sources. In the present study, the meteorological data before the 1950s were reconstructed based on the tree-ring data. Additionally, the data collected from the urban meteorological stations nearly for 60 years in the northern GKM may be affected by the enhanced urban heat island effect (He *et al.*, 2013), leading to great uncertainties in the research on the climate changes in the northern GKM during the past 300 years. If the effects of urbanization on the data collected from the meteorological stations can be removed, the research on the mechanisms of climate changes in this region can be more accurate.

Generally, climate changes may lead to the changes in forest ecosystem structure and tree species. For example, the rising temperature can enhance the competitiveness of thermophilic species and contribute to their update and succession; however, the growth of chimonophilous plants may be suppressed (Xie *et al.*, 2009; Wu *et al.*, 2014). As the precipitation



decreases, the growth of hygrophilous plants may be suppressed; even seriously, some leaves may fall or the tops may wither. This condition is conducive to the reproduction and invasion of some drought-enduring species. Climate changes can more easily change the plant's phenology and thus affect the dependency or competitiveness between species, i.e., climate changes can affect the reproduction and survival of species (Li *et al.*, 2014; Wu *et al.*, 2014). Besides, climate changes can change the plants' physiological and ecological characteristics and thereby play an important role in the formation of different species' durability, reproductive capacity and migration ability in the new system. In a word, due to climate changes, some species may exit from the original forest ecosystem while some new species invade the system, i.e., the structure and species composition of the original forest ecosystem can thus be changed (Liu *et al.*, 2015). Similarly, climate changes (for example, warming, increasing precipitation and reduced sunshine duration) not only changed the forest vegetation and the distribution pattern suitable for tree species in the northern GKM, but also may impose profound impacts on economic developments, ecology and environment in this region. Warming leads to the prolonging of plant vegetation season and the reduced occurrence probability induced by chilling damages, which will bring the newborn forestry climatic resources and can provide the possibility for the adjustment of forestry production system (Gao *et al.*, 2013). The increased precipitation means the increased probability of extreme precipitation event, which could bring adverse effects on forestry production and wetland ecosystem. Meanwhile, the decline in sunshine duration may be bad for the plants' photosynthetic efficiency and thus limited the release of the plants' potential productivity. As a consequence, how to cope with the effects of the uncertainties of climate changes on the research on the forest ecosystem in the northern GKM is a major subject in the present study, which still requires more in-depth, systematic and comprehensive studies.

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